

Infiltration and Runoff

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HyWa



Content

- 1. Introduction
- 2. Relevance in Hydrology - Runoff Generation
- 3. Formula
 - Philip
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Objectives

Understand Infiltration Processes

- Fundamental soil physics
- Know most important formulae
- Know measurement techniques
- Understand relevance for hydrology
- Be able to apply infiltration models



Infiltration Switch in Flowpaths

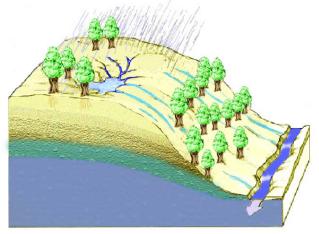
Infiltration is an important switch in hydrological flow paths, dividing precipitation into surface flows and subsurface flows.

$$Q_s = P - I \,[\mathrm{mm}]$$

mit Q_s = surface runoff in mm, *P* Precipitation, *I* Infiltration in mm.



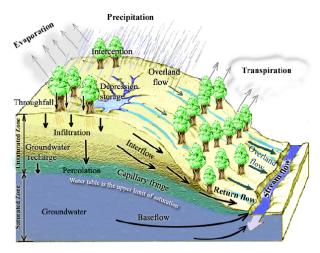
Infiltration and runoff



Tarboton, 2003

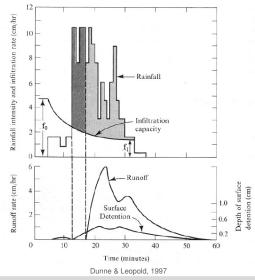


Infiltration and runoff



Tarboton, 2003



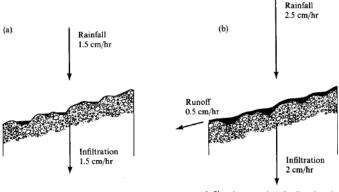


- Infiltration rate decreases
- Approaches constant value

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Infiltration and runoff



Infiltration capacity of soil = 2 cm/hr

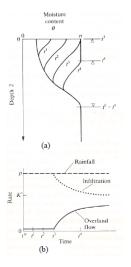
Dunne & Leopold, 1997

P > *I* infiltration and runoff *P* <= *I* infiltration, no runoff

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Infiltration and runoff





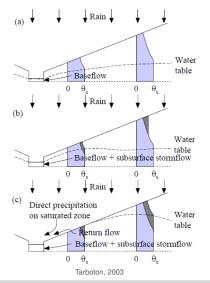
(c)

Bras, cited in Tarboton, 2003

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Infiltration and runoff

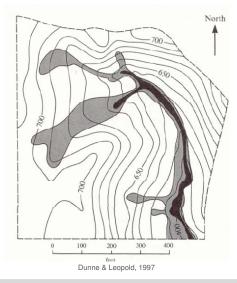


- Infiltration saturates soil
- Saturated soil produces runoff
- Floodplain turns into runoff producing area

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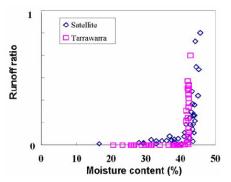
Infiltration and runoff



- Saturated area near rivers
- Runoff producing areas
- High correlation of moisture with runoff coefficient



Infiltration and runoff



Relationship between runoff ratio and soil moisture content (Woods et al., 2001, Copyright, 2001, American Geophysical Union, reproduced by permission of American Geophysical Union).

Tarboton, 2003



Transmission losses



Flood wave advancing over a dry streambed in Walnut Gulch experimental watershed where channel transmission losses are considerable (Courtesy of David Goodrich, USDA-ARS).

Goodrich, USDA cited in Tarboton, 2003



Infiltration models



Philip Infiltration Formula Infiltration Amount

Philip wrote a famous paper on the theory of infiltration in 1969: Philip J R. Theory of infiltration. Advan. Hydrosci. 5:215-96, 1969. [Division of Plant Industry, CSIRO, Canberra, Australia]

$$I = S * \sqrt{t} + a * t$$

- with *I* infiltration amount
 - S sorptivity
 - *a* hydraulic conductivity
 - t time

[*mm*] [*mm/h^{-0.5}*] [*mm/h*] [*hours*]



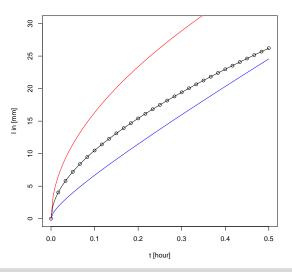
R program

The infiltration formula can be plotted using an R program.

1: S <- 30 # mm
2: a <- 8 # mm/hour
3: t <- seq(0,0.5,by=1/60.0)
4: I <- S*t^(0.5)+a*t
5: plot(t,I)</pre>



R program of infiltration amount (Philip)



 example with S=30 mm, a=10 m/hour

 low, infiltration rate, high sorption (red)

 high infiltration rate, low sorption (blue)



Philip Infiltration Formula Infiltration rate

The infiltration rate *i* in *mm* per time can be derived from the derivation of the Philip formula for infiltration amount (or vice vs. amount by integration of rate):

$$i = \frac{1}{2} * S * t^{-0.5} + a$$

- with *i* infiltration rate
 - S sorptivity
 - *a* hydraulic conductivity
 - t time

[*mm/h*] [*mm/h*^{-0.5}] [*mm/h*] [*h*]



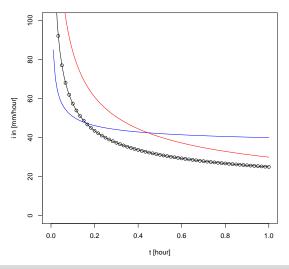
R program

The infiltration rate formula can be plotted using an R program.

1: S <- 30 # mm 2: a <- 8 # mm/hour 3: t <- seq(0,1,by=1/60.0) 4: i <- 0.5*S*t^(-0.5) + a 5: plot(t,i)



R program of infiltration rate (Philip)

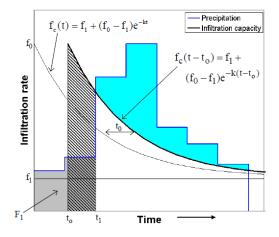


- example with S=30 mm, a=10 m/hour
- low, infiltration rate, high sorption (red)
- high infiltration rate, low sorption (blue)

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Horton equation

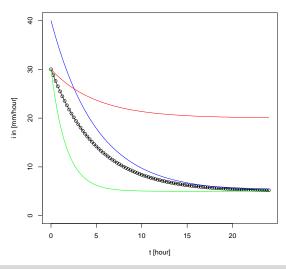


- f₁ constant infiltration rate
- *f*₀ initial infiltration rate
- k recession factor

cited in Tarboton, 2003



R program of Horton equation

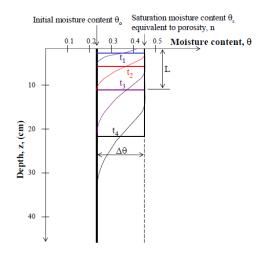


- example with $f_0 = 30 \text{ mm},$
 - $f_1 = 5$
 - mm/hour
- higher initial rate (blue)
- higher final rate (red)
- higher recession constant (green)

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Green & Ampt equation



- Assuming block sharp wetting front
- Analytical solution
- Can include ponded water

cited in Tarboton, 2003



Green & Ampt equation

At any time, t, the penetration of the infiltrating wetting front will be Z. Darcy's law can be stated as follows:

$$q = \frac{dI}{dt} = -K_s * \left[\frac{h_f - (h_s + Z)}{Z}\right]$$

where K_s is the hydraulic conductivity corresponding to the surface water content, and I(t) is the cumulative infiltration at time *t*, and is equal to $Z * (\theta_s - \theta_0)$.



Green & Ampt equation

Using this relation for I(t) to eliminate Z and performing the integration yields,

$$I = K_s * t - (h_f - h_s) * (\theta_s - \theta_0) * \log_e \left(1 - \frac{I}{(h_f - h_s) * (\theta_s - \theta_0)}\right)$$

with	I(t)	infiltration amount	[<i>cm</i>]
	Ks	hydr. conductivity	[cm/h]
	h_{f}	wetting front pressure head (negative)	ст
	h _s	water pressure at surface (ponding)	ст
	θ_s	moisture content at saturation	_
	θ_0	antecedent moisture	_

Green & Ampt equation

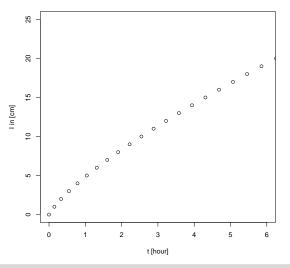
In order to solve this equation, we need to bring I(t) to one side of the equation:

$$\frac{1}{K_s} * \left[I - (h_f - h_s) * (\theta_s - \theta_0) * \log_e \left(1 - \frac{I}{(h_f - h_s) * (\theta_s - \theta_0)} \right) \right] = t$$

infiltration amount with I(t)|cm|Ks [cm/h]hydr. conductivity hf wetting front pressure head (negative) cm water pressure at surface (ponding) hs cm moisture content at saturation $\theta_{\rm S}$ antecedent moisture θ_{n}



R program of Green & Ampt equation



- implicit: insert *I* to get *t*
- plot t (result) versus l (input)
- works though!
- calculates impact of ponding

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Green & Ampt parameters

Green - Ampt infiltration parameters for various soil classes (Rawls et al., 1983). The numbers in parentheses are one standard deviation around the parameter value given.

			-	
Soil	Porosity n	Effective	Wetting	Hydraulie
Texture		porosity θ.	front soil	conductivity
		• • • •	suction	K _{1at} (cm/hr)
			head $ \Psi_f $	
			(cm)	
Sand	0.437	0.417	4.95	11.78
	(0.374-0.500)	(0.354-0.480)	(0.97-25.36)	
Loamy	0.437	0.401	6.13	2.99
sand	(0.363-0.506)	(0.329-0.473)	(1.35-27.94)	
Sandy	0.453	0.412	11.01	1.09
loam	(0.351-0.555)	(0.283-0.541)	(2.67-45.47)	
Loam	0.463	0.434	8.89	0.34
	(0.375-0.551)	(0.334-0.534)	(1.33-59.38)	
Silt loam	0.501	0.486	16.68	0.65
	(0.420-0.582)	(0.394-0.578)	(2.92-95.39)	
Sandy clay	0.398	0.330	21.85	0.15
loam	(0.332-0.464)	(0.235-0.425)	(4.42-108.0)	
Clay loam	0.464	0.309	20.88	0.1
	(0.409-0.519)	(0.279-0.501)	(4.79-91.10)	
Silty clay	0.471	0.432	27.30	0.1
loam	(0.418-0.524)	(0.347-0.517)	(5.67-	
			131.50)	
Sandy clay	0.430	0.321	23.90	0.06
	(0.370-0.490)	(0.207-0.435)	(4.08-140.2)	
Silty clay	0.479	0.423	29.22	0.05
	(0.425-0.533)	(0.334-0.512)	(6.13-139.4)	
Clay	0.475	0.385	31.63	0.03
	(0.427-0.523)	(0.269-0.501)	(6.39-156.5)	

 K_{sat} corresponds to K_s

- |ψ_f|
 corresponds
 to h_f
- θ_e
 corresponds
 to θ_s
- θ₀ needs to be known

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Summary

- Measurement methods: Guelph, double ring and suction plate
- Models: Philip, Green-Ampt, Haverkamp
- Application: Runoff modeling, recharge, pollutants